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Recovery from acidification of lochs in Galloway, south-west Scotland, UK: 1979–1998

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Abstract

The Galloway region of south-west Scotland has historically been subject to long-term deposition of acidic precipitation which has resulted in acidification of soils and surface waters and subsequent damage to aquatic ecology. Since the end of the 1970s, however, acidic deposition has decreased substantially. The general pattern is for a rapid decline in non-marine sulphate in rainwater over the period 1978–1988 followed by stable concentrations to the mid-1990s. Concentrations of nitrate and ammonium in deposition have remained constant between 1980 and 1998. Seven water quality surveys of 48 lochs in the Galloway region have been conducted between 1979 and 1998. During the first 10 years, from 1979, there was a major decline in regional sulphate concentrations in the lochs, which was expected to have produced a decline in base cations and an increase in the acid neutralising capacity. But sea-salt levels (as indicated by chloride concentrations) were approximately 25% higher in 1988 than in 1979 and thus short-term acidification due to sea-salts offset much of the long-term recovery trend expected in the lochs. During the next 10 years, however, the chloride concentrations returned to 1979 levels and the lochs showed large increases in acid neutralising capacity despite little change in sulphate concentrations. From the observed decline in sulphate deposition and concentrations of sulphate in the lochs, it appears that approximately 75% of the possible improvement in acid neutralising capacity has already occurred over the 20-year period (1979–1998). The role of acid deposition as a driving factor for change in water chemistry in the Galloway lochs is confounded by concurrent changes in other driving variables, most notably, factors related to episodic and year-to-year variations in climate. In addition to inputs of sea-salts, climate probably also influences other chemical signals such as peaks in regional nitrate concentrations and the sharp increase in dissolved organic carbon during the 1990s.

Keywords: acidification, recovery, Galloway, sulphur, nitrogen

Introduction

Acidification of surface waters in the UK uplands is linked to the emission and subsequent deposition of oxides of sulphur (S) and nitrogen (N) from the atmosphere (Battarbee *et al.*, 1990). Reconstruction of historical surface water pH, through the analysis of diatoms in lake sediment cores, shows that acidification generally occurred in the mid to late nineteenth century (DOE, 1995). The extent of acidification is determined by the sensitivity of the lake relative to the amount of acidity deposited. This sensitivity can also be related to the available pool of base cations in the catchment soils and the weathering rate of the underlying geology. In these respects, the lochs of the Galloway region of south-west Scotland represent some of the most severely acidified waters in the UK with a combination of slow

weathering granite bedrock, thin organic soils and a high deposition flux of S and N.

The first synoptic regional survey of water chemistry in the UK was carried out in Galloway in 1979 (Wright and Henriksen, 1980). The results from this survey indicate that lochs situated on the granite intrusion are extremely acidic relative to those on the surrounding sedimentary bedrock receiving similar deposition loads. A repeat survey of the same lochs in 1988 showed a 42% decrease in sulphate (SO_4) concentrations across the region resulting in a substantial increase in acid neutralising capacity (ANC) and pH in the most acidic lochs (Wright *et al.*, 1994). This was consistent with a decrease in S deposition over the same period. A further five water chemistry surveys were undertaken in the period 1988–1998 and analysis of the chemical changes from

these more recent surveys are assessed in relation to the earlier observations. This represents a unique summary of the water chemistry of the Galloway lochs from 1979–1998. The aims of this paper, therefore, are to determine the magnitude of these recent chemical changes, identify the major external drivers of change and to discuss what further potential recovery is to be expected.

Regional description

The Galloway region is a classic acid sensitive environment with highly siliceous granitic bedrock covered by thin, patchy, organic rich and generally acidic soils, which offers only limited ability to neutralise acid inputs from the atmosphere (Wright *et al.*, 1994). Decades of acid deposition and, more recently, large-scale afforestation in the region, have exacerbated the problem of soil and water acidification (Rees and Ribbens, 1995). The Galloway region also receives some of the highest rainfall in Scotland with annual average exceeding 2000 mm.

The bedrock geology of the Galloway area is characterised by Paleozoic (mainly Silurian) sedimentary rocks (shale, slate, greywacke, sandstone and conglomerate) into which granitic plutons are intruded (Daysh, 1974). The soil distribution is closely related to altitude and slope with large

areas of brown forest soils occurring below 250–300 m on very steep land and peaty podzols at higher elevations or on less steep slopes. Poorly developed alpine rankers and lithosols occur in high altitude areas (> 600 m), particularly in the Merrick mountains. Peat, often deep (> 50cm), is widely developed in this area.

To the south and west of the region, moorland communities including *Juncus effusus*-*Sphagnum recurvum*, *Carex nigra* and *Erico-Sphagnetum papillosum* are dominant in areas of dystrophic peat, peaty gleys and low base status noncalcareous gley soils (Birse and Robertson, 1976). In the 1920s, the Forestry Commission began purchasing land in the area for afforestation. Nearly 70 000 ha was purchased, of which approximately half had been planted by 1974 (Edlin, 1974). The Forestry Commission land includes much of the higher-elevation, less accessible hill country, areas that are of relatively low value for agriculture and sheep grazing. Planting included not only the native Scots pine but also a number of exotic species such as European larch, Douglas fir and Norway and Sitka spruces.

The catchments included in these surveys were selected to represent a regional distribution of acidified lochs in relation to acid deposition, forest cover, soil type and geology (Table 1 and Fig. 1).

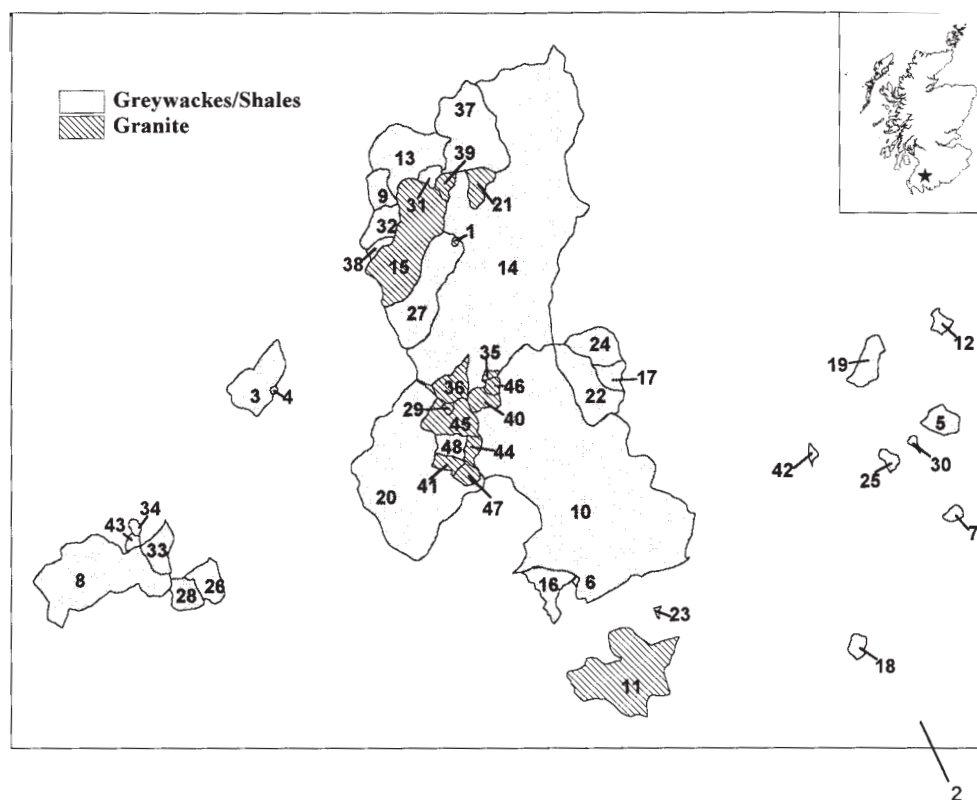


Fig. 1. Location of the study lochs in south west Scotland

Table 1. Catchment characteristics of 48 lochs in the Galloway region

Attributes		Number of lochs
Geology	Granite	14
	Greywackes	34
Soils (dominant)	Peat	5
	Peaty gley	9
	Peaty podzol	17
	Sub-alpine podzols	8
	Brown Forest Soils	9
Land Use	Forested (1988)	28
	Moorland (1988)	20
Loch:catchment area ratio	< 0.1	33
	0.1- 0.2	10
	> 0.2	5

Sampling and analysis

Surface water sampling began in 1979 and included 48 lochs in the vicinity of the three largest granitic plutons in the Galloway region (Fig. 1). These lochs were resampled in March/April of 1988, 1993, 1994, 1996, 1997 and 1998. Dip samples were collected either from the loch directly above the outflow stream or from the loch shore at about 50 cm water depth where no loch outflow existed. Samples were taken and stored in polyethylene bottles at 4°C in the dark and were filtered through 0.45 µm membrane filters prior to analysis. Samples were analysed for a full suite of chemical determinands by standard procedures (Patrick *et al.*, 1991) at the Freshwater Fisheries Laboratory, Pitlochry and the Norwegian Institute for Water Research, Oslo, for the 1979 and 1988 loch surveys. For subsequent surveys the analyses were conducted at the Macaulay Institute, Aberdeen. Strict quality control procedures between the laboratories ensured that there were no systematic differences in analytical results. pH was measured potentiometrically; calcium (Ca), magnesium (Mg), sodium (Na) and potassium (K) by atomic adsorption spectrophotometry; nitrate (NO₃), chloride (Cl), and SO₄ by ion chromatography; aluminium (Al) by automated colorimetry; dissolved organic carbon (DOC) by total carbon analyser (oxidation to carbon dioxide (CO₂) with detection by IR). ANC was calculated as Σ base cations (Ca, Mg, Na, K) minus Σ acid anions (SO₄, NO₃, Cl).

Results

DEPOSITION

An estimate of the changes in wet deposition chemistry to the Galloway region has been obtained from the Loch Dee collector (Grid Reference NX 468 779) established in 1982 as part of the UK Acid Deposition Monitoring Network (UKADMN). A longer term perspective has been obtained from wet deposition data at the Eskdalemuir collector (Grid Reference NT 234 028), also as part of the UKADMN, with records dating back to 1973. This collector is located approximately 80 km to the east and 25 km to the north of the Loch Dee site.

The deposition data from the Eskdalemuir collector follows two different protocols; mean annual concentrations are calculated from both mean daily concentrations and mean weekly concentrations, the latter forming the basis of the UKADMN data (RGAR, 1983, 1987, 1993, 1997). Comparison of the wet deposited non-marine SO₄ concentration from the two methods at Eskdalemuir (Fig. 2) shows no difference. Within the paper the non-marine component of sulphate (*nmSO*₄) was determined by:

$$nmSO_4 = SO_4 - (R * Cl)$$

where *nmSO*₄ is the non-marine concentration of SO₄, and *R* is the ratio of that ion to *Cl* in seawater (0.104 for SO₄ where concentrations are in µeq l⁻¹). The Loch Dee data is based on mean weekly concentrations and demonstrates greater annual variability in concentrations but the same overall trend as at Eskdalemuir. The general pattern is for a rapid decline in *nmSO*₄ over the period 1978–1988 approximating to a > 50% reduction, followed by more stable concentrations during the 1980s and early 1990s, and a subsequent reduction (10% of 1978 values) until 1998. Concentrations of NO₃ and ammonium (NH₄) at both

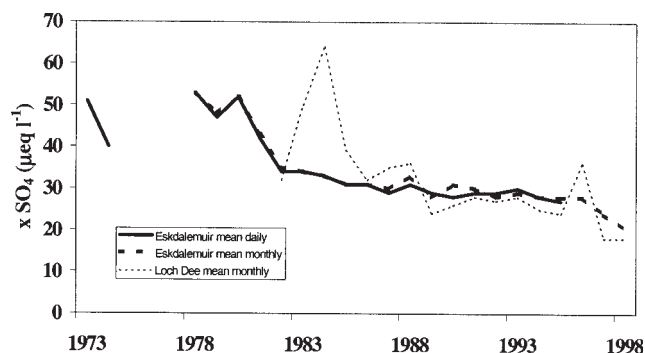


Fig. 2. Wet deposition of non-marine SO₄ at Eskdalemuir and Loch Dee

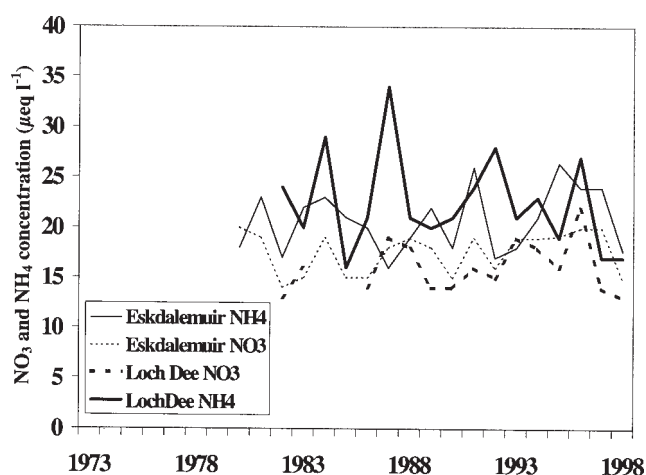


Fig. 3. Concentrations of NO_3 and NH_4 in wet deposition at Eskdalemuir and Loch Dee

Eskdalemuir and Loch Dee have remained constant between 1980 and 1998 (Fig. 3).

Similarly, concentrations of Ca and Mg have remained relatively low and constant between 1980 and 1998 (Fig. 4). Chloride concentrations, however, have been extremely variable at Loch Dee reflecting the proximity of this site to the coast.

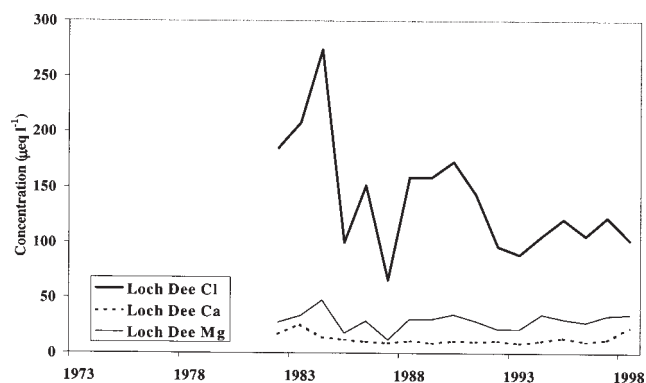


Fig. 4. Base cation and Cl concentrations in wet deposition at Loch Dee, 1980-1998

HYDROCHEMISTRY

The distribution of changes in the concentration of each ion between each of the surveys can be illustrated with ranked-slope plots for each interval between surveys (1979-1988, 1988-1998 and 1979-1998). The change in each ion for each interval is expressed as a rate of change for the interval in question by dividing the observed change in the

concentration of each ion (in $\mu\text{eq l}^{-1}$) by the number of years in the interval. These slopes are then sorted from lowest to highest and plotted against the % rank of each value. The resulting ranked-slope plots can be used to determine the percentage of sites that had slopes greater or less than some particular value.

Concentrations of SO_4 in all of the 48 lochs declined sharply during the first 10 years 1979-1988 (Figs. 5 and 6). Median change for the 48 lochs was $-6 \mu\text{eq l}^{-1}$ per year (Table 2). During the next 10 years, 1988-1998, however, SO_4 concentrations remained relatively constant with about equal numbers of lochs showing positive and negative changes (Fig. 7). This pattern of SO_4 concentrations in the lochs closely reflects the trend in SO_4 deposition in the region; the major decline in SO_4 deposition occurred during the early 1980s (Fig. 2).

Concentrations of NO_3 in the lochs, on the other hand, show no systematic regional pattern during the 20 years (Fig. 5). Concentrations were somewhat lower in 1988 relative to 1979, but then increased again during the period 1988-1998 (Figs. 6 and 7). In addition, deposition of inorganic N in the region has not changed systematically over this period. Samples collected in spring 1996 had exceptionally high concentrations of NO_3 (Fig. 5).

Concentrations of sea-salt-derived ions, particularly Cl, in the lochs varied greatly, but did not show any major trend over time for the region (Fig. 5). Levels were relatively low in the first survey, higher in 1988 and then lower again during the late 1990s (Table 2). Linear changes are thus

Table 2. Median regional concentration of major ions in 1998 and slopes of changes for the periods 1979-1988, 1988-1998 and 1979-1998. All data as $\mu\text{eq l}^{-1}$, all changes presented as $\mu\text{eq l}^{-1}$ per year.

	Median regional chemistry 1998	Change 1979- 1988	Change 1988- 1998	Change 1979- 1998
Ca	49.9	-1.35	0.34	-0.24
Mg	49.9	-0.73	-0.07	-0.11
Na	171.9	3.43	-1.83	1.26
K	7.6	-0.23	0.10	-0.04
NO_3	8.8	-0.76	0.68	0.07
SO_4	59.1	-5.78	-0.39	-2.41
Cl	180.2	5.64	-5.50	0.22
SBC	281.3	0.50	-1.42	1.07
SAA	248.1	-0.74	-6.01	-1.97
ANC	33.2	1.18	4.50	3.26

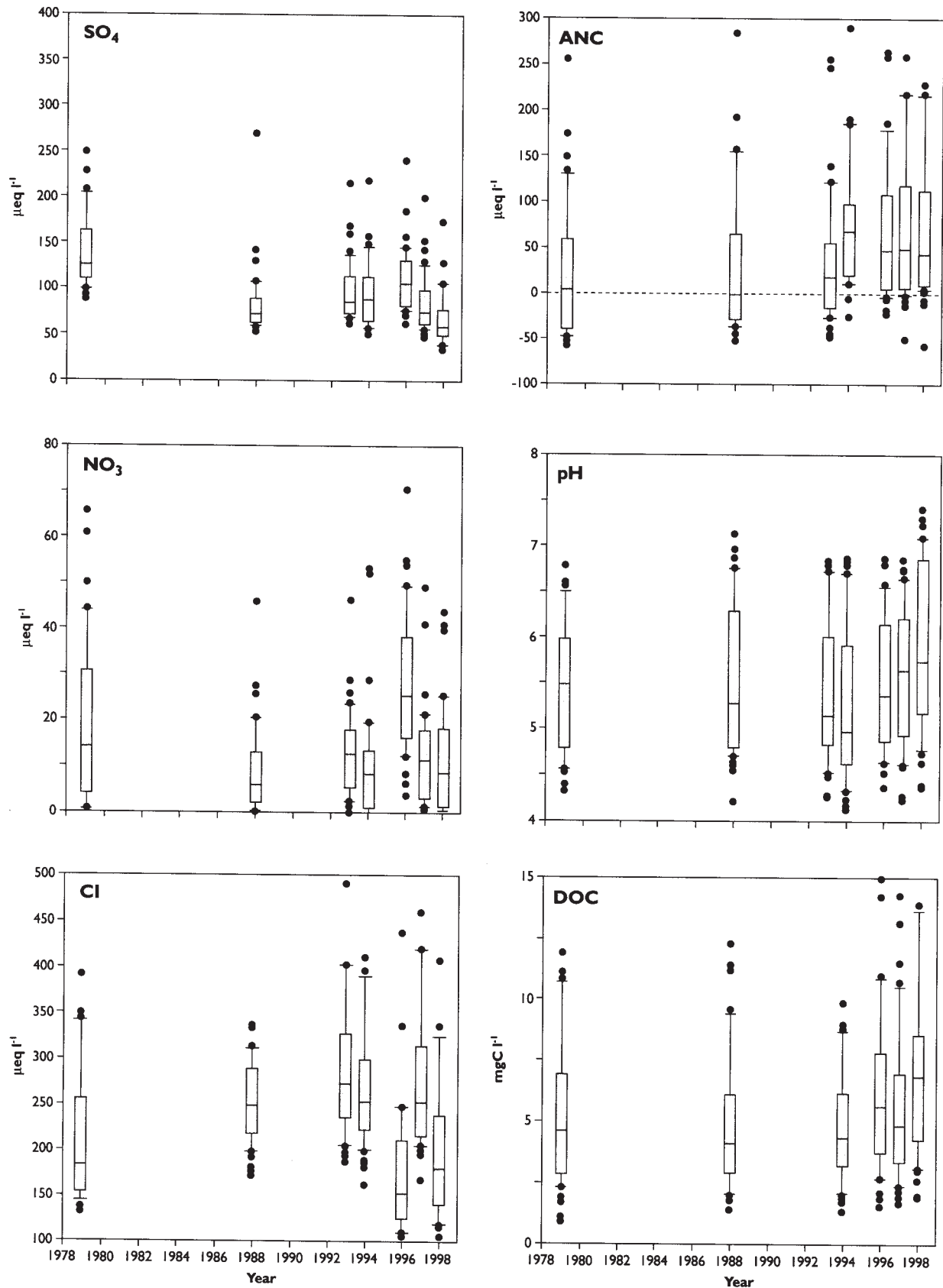


Fig. 5. Box and whisker plots of selected determinands (SO_4 , NO_3 , C, ANC, pH and DOC) for the seven regional surveys in Galloway, since 1979. All data as $\mu\text{eq l}^{-1}$ except pH, and DOC (mgC l^{-1}). The box represents the 25 and 75 percentile and the whisker is the 10th and 90th percentile, outliers are also shown.

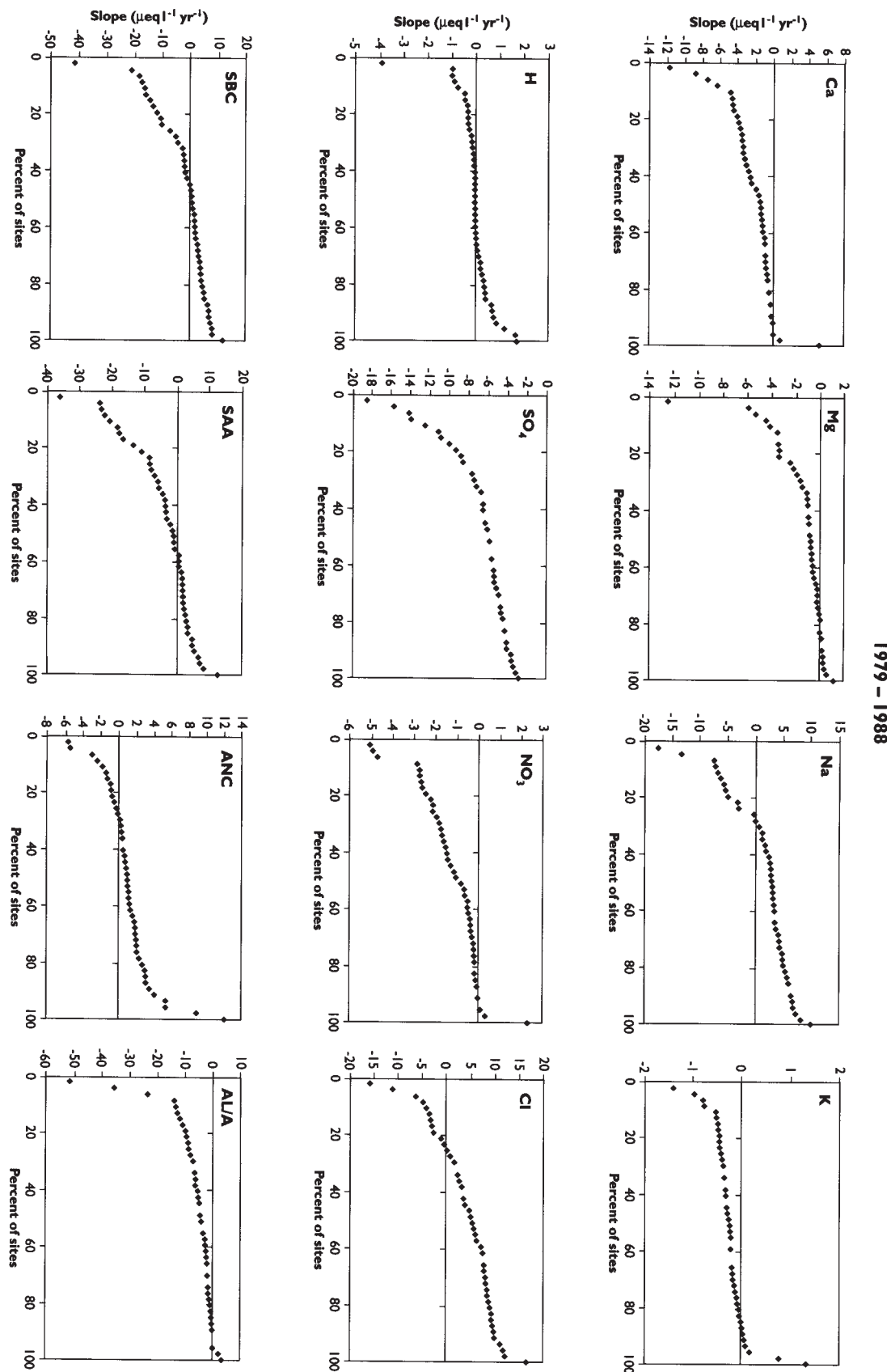


Fig. 6. Percentage of sites showing positive or negative trends for water quality variables in the period 1979-1988. All data as $\mu\text{eq l}^{-1}$

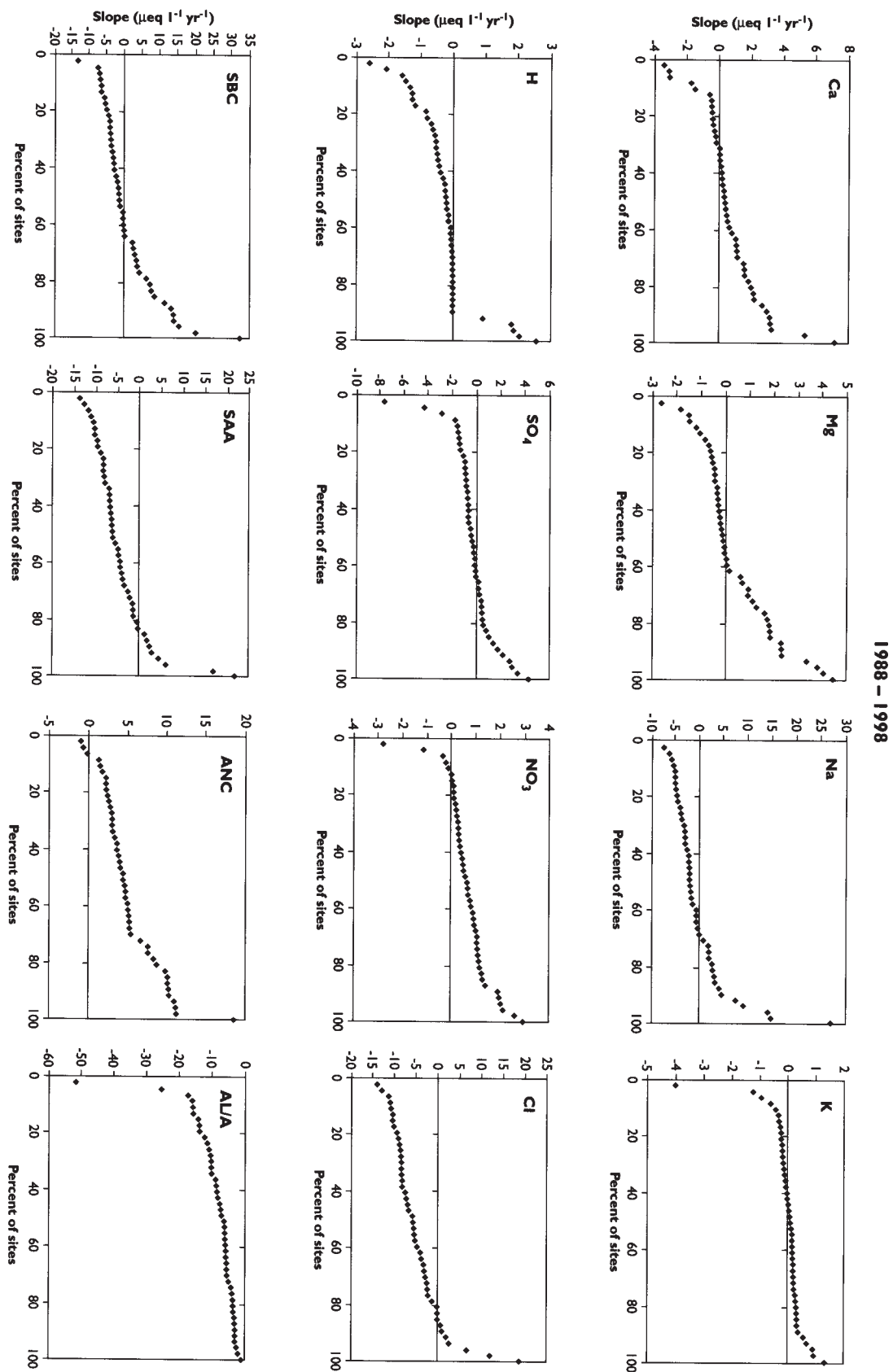


Fig. 7. Percentage of sites showing positive or negative trends for water quality variables in the period 1988–1998. All data as µeq l⁻¹.

positive in the first 10 years (Fig. 6) and negative during the second 10 years (Fig. 7). In this instance, no significant regional trend was observed during the 20 year record (Fig. 8).

As a consequence of the changes in SO_4 , NO_3 and Cl, the median sum of strong acid anions (SAA) decreased by about $1 \mu\text{eq l}^{-1}$ per year over the period 1979 to 1988; the large decrease in SO_4 was offset by an increase in Cl. Over the second 10 year period (1988–1998), the median SAA decreased further by $6 \mu\text{eq l}^{-1}$ per year, largely due to lower Cl concentrations (Table 2). For the entire 20 years the median change in SAA was $-2 \mu\text{eq l}^{-1}$. All but four of the lochs showed negative change in SAA over the entire period (Fig. 8).

Charge balance necessitates that decreases in SAA are compensated by either a decrease in base cations or an increase in ANC. Median change in sum of base cations was $+1 \mu\text{eq l}^{-1}$ per year over the entire 20 years (Table 2). Trend analyses of Ca and SO_4 are very similar. Loch data from the first 10 years showed a decline in Ca (median slope of $-1.4 \mu\text{eq l}^{-1}$ per year), with little change during the next 10 years. Trends in Mg and K are similar to those for Ca.

During the first and second 10 year period, increases in ANC were concurrent with decreases in SAA concentrations, whereby the median slope of ANC in the first 10 years equalled $+1 \mu\text{eq l}^{-1}$ per year and $+4 \mu\text{eq l}^{-1}$ per year during the second 10 year period (Figs. 6, 7, and 8; Table 2). The increases in ANC were due to decreased concentrations of hydrogen (H) ion and Al. Dissolved organic carbon (DOC) appears to have increased from 1994 to 1998.

Discussion

Monitoring of surface water chemistry began in earnest in the UK in the late 1970s following the recognition that surface water acidification presented a problem. The approach to monitoring has taken two directions: systematic and routine sampling of individual sites at monthly intervals (or less) and synoptic surveys sampling surface standing waters in a region on an annual frequency (or more). Determination of chemical trends from the former relies upon detecting a signal from noise introduced by the short-term chemical response to rainfall input and seasonality. In the latter, the signal must be detected from noise introduced by different physical characteristics of a large population of lakes. In practice, a balance between both approaches provides the most robust results and interpretation.

The data presented come from a regional survey of 48 lochs conducted seven times during a 20-year period. The general pattern of change identified by these spatially

extensive but temporally sparse data, mirrors changes in individual lochs sampled more frequently over the same time period by Harriman *et al.* (1995, 2001) and NEG-TAP (2001). Temporal changes in loch chemistry identified in this study, therefore, provide a mechanism to examine the regional response of both highly acidified as well as more moderately acidified waters. For example, the data of Harriman *et al.* (2001) for Loch Enoch show large decreases in SO_4 concentrations during the 1980s with smaller changes during the 1990s, which are concomitant with deposition trends highlighted in Fig 2. This pattern of change is also accompanied by small decreases in concentrations of base cations and an increase in ANC (Harriman *et al.*, 2001). The seven regional surveys conducted over 20 years, superimposed on detailed chemistry from high-frequency monitoring at selected sites, thus appear to capture adequately the major trends in water chemistry in Galloway lochs.

Surface waters in the Galloway region are strongly influenced by inputs of sea-salts from the atmosphere. This is generally the case over large areas of the UK (Monteith and Evans, 2000). Inputs of sea-salts occur in conjunction with storms and since Cl is very mobile in soils, Cl concentrations in streams and lochs show large variation over time. The cations in the sea-salts, mainly Na and Mg, are not as mobile in soils as Cl since both cations participate in cation exchange reactions. In acidified soils, such as those typical of the Galloway region, the incoming Na and Mg exchange in part for inorganic Al and H, resulting in decreased ANC of the run-off. This episodic acidification due to sea-salt deposition has been documented in studies from Loch Dee in Galloway (Langan, 1987), Norway (Hindar *et al.*, 1994) and Maine, USA, (Heath *et al.*, 1992) and was demonstrated experimentally at Sogndal, Norway (Wright *et al.*, 1988).

Episodic input of sea-salts must be considered in case they obscure ongoing trends of acidification, as is apparently the case in this 20-year record from the Galloway lochs. It is especially important, therefore, to ensure that low frequency regional survey data is consistent with high frequency data collected from sites in the same region. The major decline in SO_4 concentrations in the lochs during the first ten years should have produced declines in base cations and increases in ANC. But the sea-salt levels (as shown by Cl concentrations) were about 25% higher in 1988 than in 1979 and thus short-term acidification due to sea-salts offset much of the long-term recovery expected in the lochs. During the next 10 years, however, Cl concentrations decreased back to 1979 levels and the lochs show large increases in ANC despite little change in SO_4 concentrations. The trend data indicate that year-to-year variations in sea-

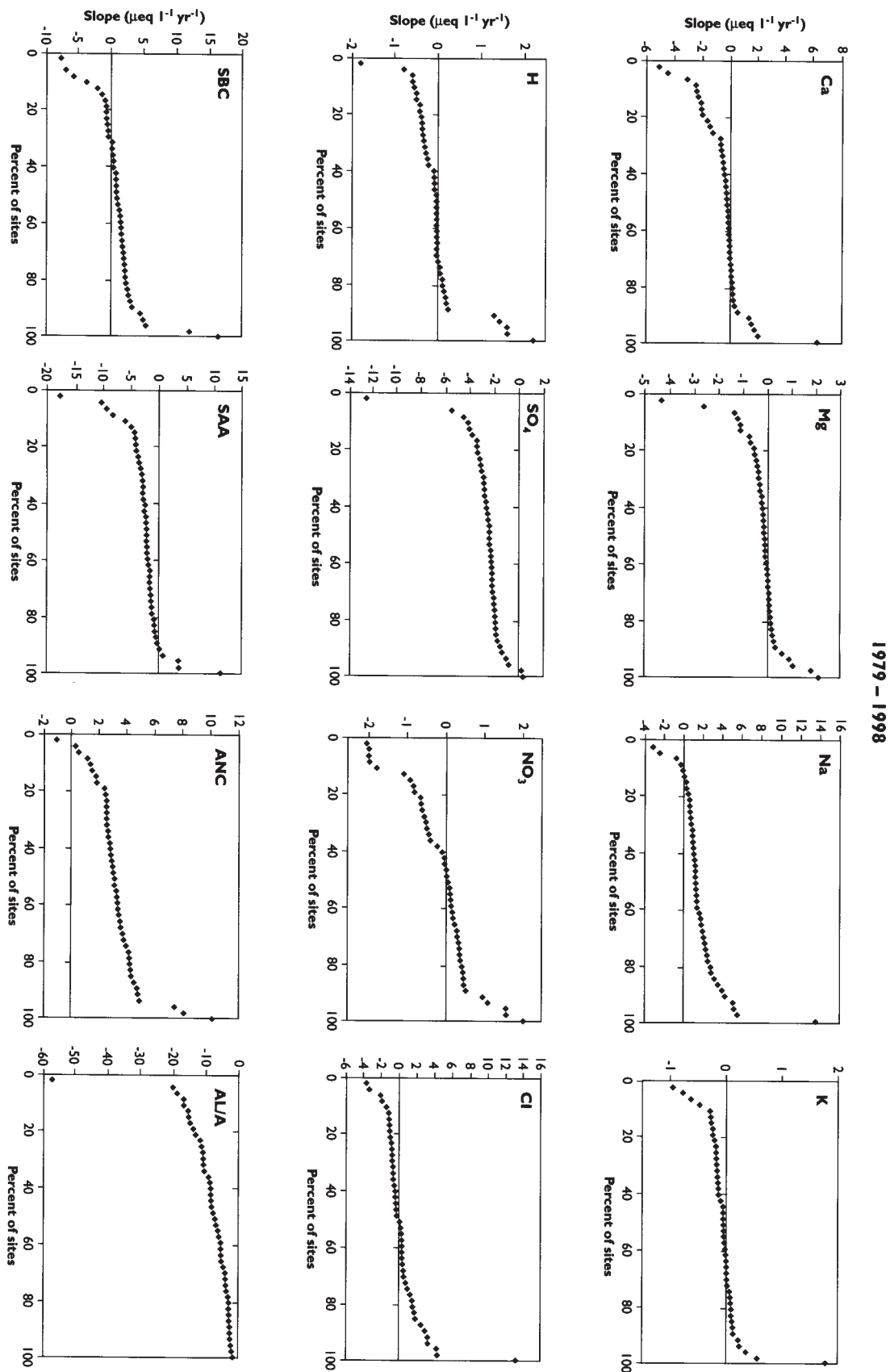


Fig. 8. Percentage of sites showing positive or negative trends for water quality variables in the period 1979-1998. All data as µeq l⁻¹.

salt deposition caused a 5–10 year delay in the recovery in ANC in the lochs. This is a similar result to that of Monteith and Evans (2000) in their analysis of a 10-year record of water chemistry trends in 22 lakes and streams in the UK Acid Waters Monitoring Network (AWMN). They showed that variable inputs of sea-salts causes ‘noise’ in the concentrations of major ions such that long-term trends may be partly masked.

Nitrate concentrations in streamwaters have been shown to be related to the amount of N deposition with sites receiving between 10 and 25 kg N ha⁻¹ yr⁻¹ showing a variable response, such that some sites show elevated levels, whilst others do not. Sites with deposition inputs above 25 kg N ha⁻¹, however, have shown elevated NO₃ outputs (Dise and Wright, 1995). For Galloway, present day deposition of N is higher than the 25 kg N ha⁻¹ threshold and some evidence of N enrichment and potential elevated N release should be expected. The current 20-year records suggest that there has been no regional increase in N saturation, at least under the criterion of increased leaching of inorganic N from the soil (Aber *et al.*, 1989). The lochs and their catchments retain 80–100% (Helliwell *et al.*, 2001) of inorganic N deposition and have done so for the 20-year period. The number of lochs with increasing NO₃ concentrations is equal to those with decreasing trends over the 20 year period. If N saturation is occurring, the rate of breakthrough is very slow (i.e. very long time lag). Helliwell *et al.* (2001), however, highlight that N leaching was evident from forested catchments in Galloway with soil C/N ratios below 20, in line with that predicted for European forest soils (Gundersen, 1995; Gundersen *et al.*, 1998).

The exceptionally high NO₃ concentrations in 1996 are apparently a larger regional phenomenon, as such a peak is also reported from the 10-year records of the AWMN (Monteith *et al.*, 2000). The cause of this peak is probably related to climate and Monteith *et al.* (2000) suggest that unusually cold winters give rise to higher NO₃ concentrations in surface waters in the spring.

From the decline in SO₄ deposition and concentrations of SO₄ in the lochs, it appears that about 75% of the possible improvement in surface water quality has already occurred over the 20-year period 1979–1998. The median non-marine SO₄ concentration in the lochs declined by 67 µeq l⁻¹ from about 108 µeq l⁻¹ in 1979 and to about 41 µeq l⁻¹ in 1998. The marine fraction of SO₄ amounts to about 20 µeq l⁻¹. Median non-marine SBC declined by about 23 µeq l⁻¹ from about 115 to 92 µeq l⁻¹, while median ANC increased by about 40 µeq l⁻¹ from about 3 to about 43 µeq l⁻¹. The decline in SO₄, therefore, was compensated one third by decreased SBC and two thirds by increased ANC. From these changes, a rough estimate of the maximum future recovery in ANC

can be obtained. If in the future non-marine SO₄ deposition were to decrease to zero, the concentrations in the lochs would decrease by a further 20 µeq l⁻¹ to median value of about 20 µeq l⁻¹ (an estimate of median background non-marine SO₄ concentration). This would be compensated by about a 7 µeq l⁻¹ decrease in SBC and a 13 µeq l⁻¹ increase in ANC. This rough estimate indicates, therefore, that of the total improvement in ANC (median change of 53 µeq l⁻¹), about 75% (median change 1979 to 1998 40 µeq l⁻¹) has already occurred.

This estimate assumes no long-term changes in NO₃ concentrations. Nitrate in the future could go either up, down, or stay the same. Continued inputs of inorganic N from deposition should, over the long term, result in N saturation with increased concentrations of NO₃ in run-off (Aber *et al.*, 1989). The 20-year record from Galloway lochs indicates, however, that this process is very slow, as there has been no significant increase in NO₃ concentrations in the lochs over the period 1979–98 despite 20 years of high N deposition in the area. On the other hand, NO₃ concentrations in the future could decrease due to decreased deposition of inorganic N. If the emission reduction requirements of the Gothenburg protocol are indeed implemented, then N deposition in Europe should decrease by over 50% by the year 2010 relative to the base year 1990 (expected reductions: SO₄ 63%; NO_x 41%; NH_y 17%). But in this case, since median NO₃ concentration in 1998 was only about 10 µeq l⁻¹, future decreases in concentrations could only cause a few µeq l⁻¹ increase in ANC. The potential impact of land use and management on terrestrial N cycling, in particular commercial afforestation, is also unclear (Ferrier *et al.*, 1995).

The increase in DOC concentrations during the 1990s in the Galloway lochs is part of a general trend reported from many areas of northwestern Europe. The 10-year record from the UKAWMN shows striking increases in DOC in most sites (Monteith and Evans, 2000), as do similar data from waters in Norway (Skjelkvåle *et al.*, 1998). Here, again, climate is probably the driving factor as the increases have occurred simultaneously in waters over a very large area and in both acidified and non-acidified systems (Monteith and Evans, 2000). Increased DOC will result in increased organic anion availability which will buffer pH and increase ANC. Thus, the pH levels in the lochs in Galloway have recovered less than expected from the increase in ANC.

In summary, it is clear that the lochs in Galloway have shown substantial recovery in response to a long-term decline in the anthropogenic deposition of primarily S and, to a lesser degree, N. The regional pattern of recovery is consistent with that observed for specific lochs where high frequency monitoring has been carried out. The role of acid

deposition as a driving factor for changes in water chemistry in the Galloway lochs, however, is confounded by concurrent changes in other driving variables, most notably factors related to variations in climate. Sea-salt deposition and other climatic influences, are regionally important in this area of Scotland and strongly influence the observed hydrochemical response, especially N and C dynamics.

Acknowledgements

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